



Terrestrial planet formation constrained by Mars and the structure of the asteroid belt

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ABSTRACT

Reproducing the large Earth/Mars mass ratio requires a strong mass depletion in solids within the protoplanetary disc between 1 and 3 au. The Grand Tack model invokes a specific migration history of the giant planets to remove most of the mass initially beyond 1 au and to dynamically excite the asteroid belt. However, one could also invoke a steep density gradient created by inward drift and pile-up of small particles induced by gas drag, as has been proposed to explain the formation of close-in super-Earths. Here we show that the asteroid belt's orbital excitation provides a crucial constraint against this scenario for the Solar system. We performed a series of simulations of terrestrial planet formation and asteroid belt evolution starting from discs of planetesimals and planetary embryos with various radial density gradients and including Jupiter and Saturn on nearly circular and coplanar orbits. Discs with shallow density gradients reproduce the dynamical excitation of the asteroid belt by gravitational self-stirring but form Mars analogues significantly more massive than the real planet. In contrast, a disc with a surface density gradient proportional to $r^{-5.5}$ reproduces the Earth/Mars mass ratio but leaves the asteroid belt in a dynamical state that is far colder than the real belt. We conclude that no disc profile can simultaneously explain the structure of the terrestrial planets and asteroid belt. The asteroid belt must have been depleted and dynamically excited by a different mechanism such as, for instance, in the Grand Tack scenario.

Key words: methods: numerical – planets and satellites: formation.

1 INTRODUCTION

One goal of the field of planet formation is to reproduce the Solar system using numerical simulations (for recent reviews, see Morbidelli et al. 2012; Raymond et al. 2014). Most studies have not achieved this goal. A number of fundamental properties of the Solar system are difficult to replicate. In this paper we focus on two of these key constraints: Mars' small mass and the structure of the asteroid belt (orbital distribution, low total mass).

The classical scenario of terrestrial planet formation suffers from the so-called 'Mars problem' (e.g. Chambers 2014). Assuming that planets accrete from a disc of rocky planetesimals and planetary embryos that stretches continuously from ~ 0.3 – 0.7 au to about 4–5 au, simulations consistently reproduce the masses and orbits of Venus and Earth (Wetherill 1978, 1986; Chambers & Wetherill 1998; Agnor, Canup & Levison 1999; Chambers 2001; Raymond,

Quinn & Lunine 2004, 2006, 2007a; O'Brien, Morbidelli & Levison 2006; Morishima et al. 2008; Raymond et al. 2009; Morishima, Stadel & Moore 2010; Izidoro et al. 2013; Lykawka & Ito 2013; Fischer & Ciesla 2014). However, planets in Mars' vicinity are far larger than the actual planet. Several solutions to this problem have been proposed, each invoking a depletion of solids in the Mars region linked to either the properties of the protoplanetary disc (Jin et al. 2008; Hansen 2009; Izidoro et al. 2014), perturbations from the eccentric giant planets (Raymond et al. 2009; Morishima et al. 2010; Lykawka & Ito 2013), or a combination of both (Nagasawa, Lin & Thommes 2005; Thommes, Nagasawa & Lin 2008). Most of these models are either not self-consistent or are simply inconsistent with our current understanding of the global evolution of the Solar system (for a discussion see Morbidelli et al. 2012).

To date the most successful model – called the *Grand Tack* – invokes a truncation of the disc of terrestrial building blocks during the inward-then-outward migration of Jupiter in the gaseous protoplanetary disc (Pierens & Raymond 2011; Walsh et al. 2011; Jacobson & Morbidelli 2014; O'Brien et al. 2014; Raymond &

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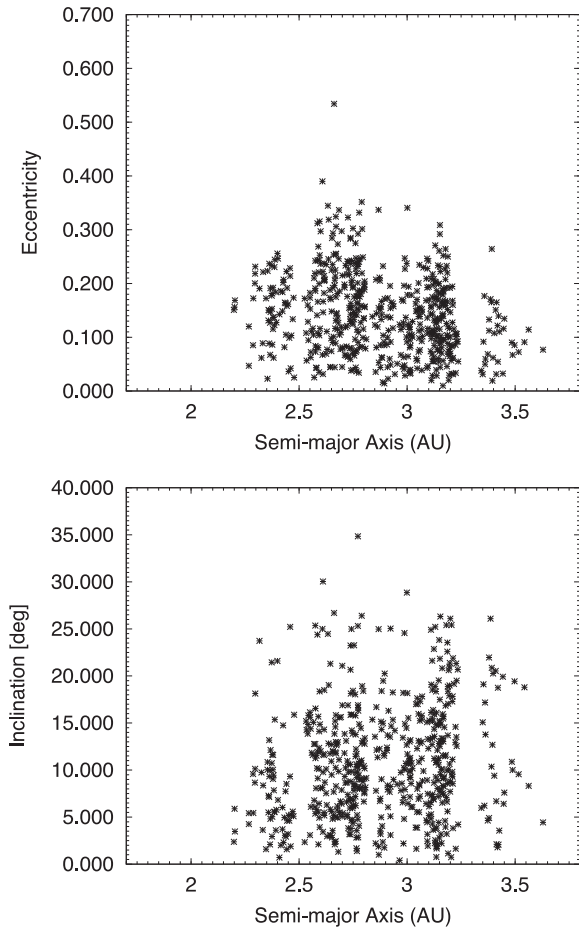


Figure 1. Orbital distribution of the real population of asteroids with absolute magnitude $H < 9.7$, which corresponds to diameter of about 50 km. The upper plot shows semimajor axis versus eccentricity. The lower plot shows semimajor axis versus inclination.

Morbidelli 2014). If the terrestrial planets formed from a truncated disc, then Mars’ small mass is the result of an ‘edge’ effect, as this planet was scattered out of the dense annulus that formed Earth and Venus (Wetherill 1978; Morishima et al. 2008; Hansen 2009).

A second model constraint comes from the asteroid belt, specifically the distribution of orbital inclinations. The present-day main belt has a broad inclination distribution, spanning continuously from $i = 0^\circ$ to 20° . Fig. 1 shows the orbital distribution of the real population of asteroids with diameter larger than about 50 km. The inclinations of most of these objects are far larger than would be expected from formation in a dissipative protoplanetary disc (Lecar & Franklin 1973).

At least four models have been proposed for explaining the belt’s depletion and excitation (for a detailed discussion see a review by Morbidelli et al. 2015): sweeping secular resonances¹ driven by the depletion of the nebula (Lecar & Franklin 1973, 1997), scattering of Earth mass objects by the forming Jupiter (Safronov 1979; Ip

1987; Petit, Morbidelli & Valsecchi 1999), scattering by protoplanetary embryos embedded in the belt (Wetherill 1992; Chambers & Wetherill 2001; Petit, Morbidelli & Chambers 2001) and the inward scattering of planetesimals during the outward migration of Jupiter and Saturn (i.e. the Grand Tack model: Walsh et al. 2011, 2012). Two of these models have been discarded because they do not reproduce the observed belt structure: sweeping of secular resonances and planets scattered by Jupiter (see O’Brien, Morbidelli & Bottke 2007; Morbidelli et al. 2015). For example, O’Brien et al. (2007) showed that sweeping secular resonances are incapable of giving the observed orbital excitation of the asteroid belt unless the time-scale for nebular gas depletion is much longer (~ 20 Myr) than the values derived from current observations (1–10 Myr).

Within the classical scenario of terrestrial planet formation, the most successful model to excite and deplete the belt invokes the existence of planetary embryos in the primitive asteroid belt (Chambers & Wetherill 2001; Petit et al. 2001; O’Brien et al. 2007). These putative planetary embryos depleted the main belt and excited the eccentricities and inclinations of the orbits of surviving asteroids. In this scenario, the asteroid belt is initially massive (roughly 1–2 Earth masses, as numerous massive planetary embryos are needed to excite the belt), while it is mass-depleted in the end. The initial presence of a significant amount of mass in the asteroid belt, though, has the drawback of producing a too massive planet at the location of Mars. In addition, all planetary embryos need to be removed from the belt on a 100 Myr time-scale; if planetary embryos are removed on a significantly longer time-scale they carve a ‘hole’ in the observed asteroid distribution (Petit et al. 1999; Raymond et al. 2009).

Unlike the classical scenario, the orbital excitation and mass depletion of the asteroid belt in the Grand Tack model are both essentially created during the inward-then-outward migration of Jupiter and Saturn (Walsh et al. 2011; Jacobson & Walsh 2015). In this model, the giant planets crossed the asteroid belt twice. First, during their inward migration stage, the gas giants compress the distribution of planetary embryos and planetesimals inside Jupiter’s orbit into a narrow disc around 1 au. A fraction of these objects is also scattered outwards. Secondly, during the outward migration phase, the giant planets scatter inwards a fraction of planetesimals beyond 2–3 au enough to repopulate the asteroid belt region, with a dynamically excited population of small bodies carrying altogether a small total mass.

An alternative to the Grand Tack scenario to produce the confined disc and a mass deficient asteroid belt could be invoked so that a lot of solid material drifted to within 1 au by gas drag, leaving the region beyond 1 au substantially depleted in mass. This idea is very appealing in a broad context of planet formation. This is because it is often invoked to produce a large pile-up of mass in the inner disc to explain the formation of close-in super-Earths (e.g. Boley & Ford 2013; Chatterjee & Tan 2014). Moreover, it could also be consistent with modern ideas on planetesimal formation and planetary growth based on the drift and accretion of pebbles (Lambrechts & Johansen 2014; Lambrechts, Johansen & Morbidelli 2014; Johansen et al. 2015). Particles drifting towards the star can produce in principle discs of solids of any radial gradient in the resulting mass distribution. Therefore, the goal of this paper is to test whether any of these gradients could explain at the same time the small mass of Mars and the properties of the asteroid belt (mass deficit and inclination excitation). In other words, can we match these constraints *without* invoking a dramatic event within the inner Solar system (such as a Grand Tack)?

¹ A secular resonance occurs when the period of the nodal (longitude of the ascending node) or apsidal (longitude of perihelion) precession of the orbit of a small body becomes equal to those of one of the giant planets. Equal apsidal frequencies result in a pumping of the orbital eccentricity of the smaller object, while a match between nodal frequencies tend to increase the orbital inclination of the smaller body.

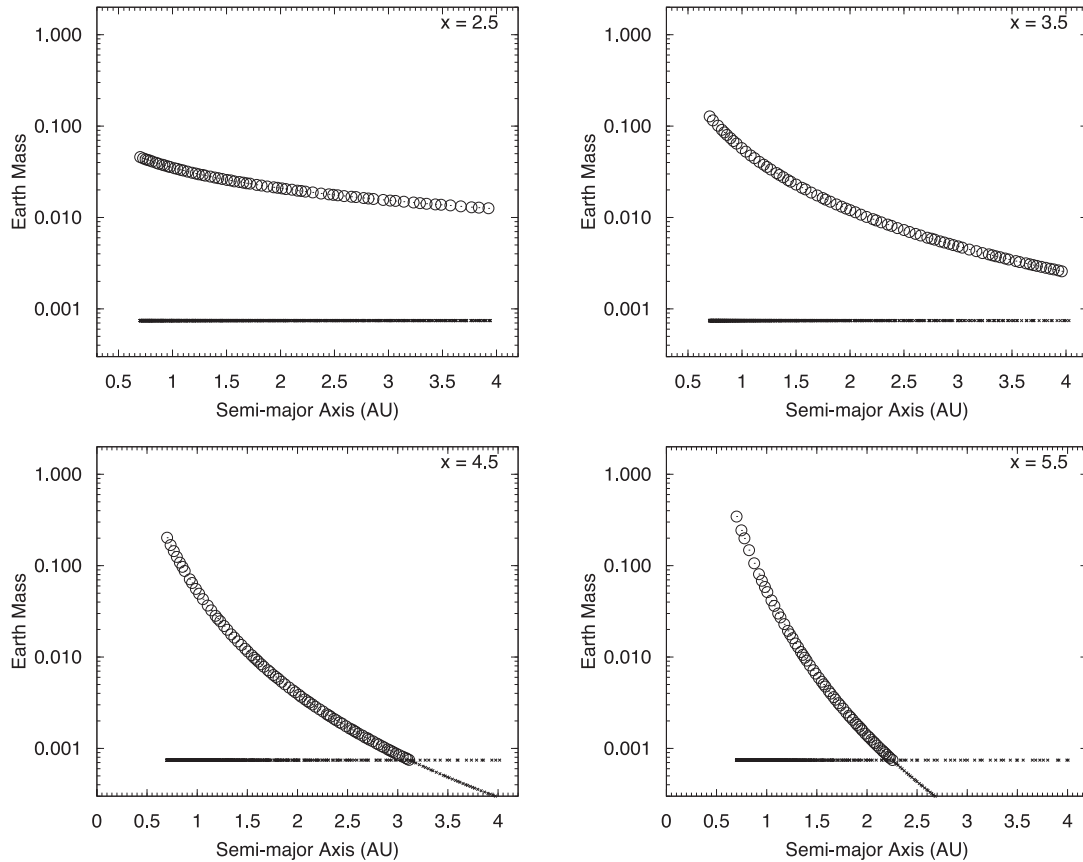


Figure 2. Initial conditions of our simulations. Each panel shows the distribution of planetary embryos and planetesimals generated within a given disc profile defined by the surface density slope x . Planetary embryos are marked as open circles and planetesimals are objects with masses smaller than 0.000 75 Earth masses.

This paper is laid out as follows. We describe the details of our model in Section 2. In Section 3 we present and analyse our results. In Section 4 we discuss our results and present our conclusions.

2 THE MODEL AND NUMERICAL SIMULATIONS

The simulations presented in this paper fit in the context of the classical scenario of terrestrial planet formation. We perform simulations starting from discs with a wide range of surface density profiles. We do not interpret these disc profiles as reflecting the properties of the primordial gaseous disc, as they are not consistent with viscous disc models (e.g. Raymond & Cossou 2014). Rather, we assume that the distribution of solids has been sculpted by other processes such as aerodynamic drift (Adachi, Hayashi & Nakazawa 1976; Weidenschilling 1980). In an infinite disc, the drift of particles would create a steady flow and no steep radial gradient of mass. But, if Jupiter formed early it could have acted as a barrier to inward drifting pebbles or planetesimals (or even planetary embryos; Izidoro et al. 2015). Thus, if a pressure bump existed in the terrestrial planet formation zone to stop the inward drift (e.g. Haghighipour & Boss 2003), these two effects together could potentially have led to the creation of a very steep density profile in the terrestrial planet region and throughout the asteroid belt.

We tested discs with surface density profiles given by $\Sigma_1 r^{-x}$, where $x = 2.5, 3.5, 4.5$ or 5.5 . Σ_1 is the solid surface density at 1 au. Our discs extend from 0.7 to 4 au. These profiles are much steeper than previous simulations, which were almost always

limited to $x = 1.5$ and 1 (except for Raymond, Quinn & Lunine 2005; Kokubo, Kominami & Ida 2006; Izidoro et al. 2014). We adjusted Σ_1 to fix the total mass in the disc between 0.7 and 4 au at $2.5 M_\oplus$, comparable to the sum of the masses of the terrestrial planets.

The disc is divided into populations of planetesimals (30–40 per cent by mass) and planetary embryos (60–70 per cent). Planetary embryos are assumed to have formed by oligarchic growth and are thus randomly spaced by 5–10 mutual Hill radii (Kokubo & Ida 1998, 2000). Individual planetesimals were given masses of 0.000 75 Earth masses or smaller. Planetesimals are assumed to interact gravitationally only with the protoplanetary embryos, giant planets and the star, but not with each other. The masses of the planetary embryos scale as $M \sim r^{3(2-x)/2} \Delta^{3/2}$ (Kokubo & Ida 2002; Raymond et al. 2005, 2009) where Δ is the number of mutual Hill radii separating adjacent orbits. This amounts to roughly 80 planetary embryos and 1000 planetesimals. Fig. 2 shows the initial conditions of our simulations. The initial individual embryo mass at a given orbital distance can vary between different discs by up to a factor of ~ 10 . Steeper discs (with higher values of x) have more massive planetary embryos in the inner parts of the disc, although they never exceed $0.3 M_\oplus$. Steep discs also have smaller planetary embryos farther out. In our steepest discs the embryo mass actually drops below the planetesimal mass in the asteroid region (and they become non-self-gravitating bodies). The lack of gravitational interactions among non-self-gravitating bodies (necessary in order to keep the computational times manageable) has important implications for this study. We will pay special attention to this issue

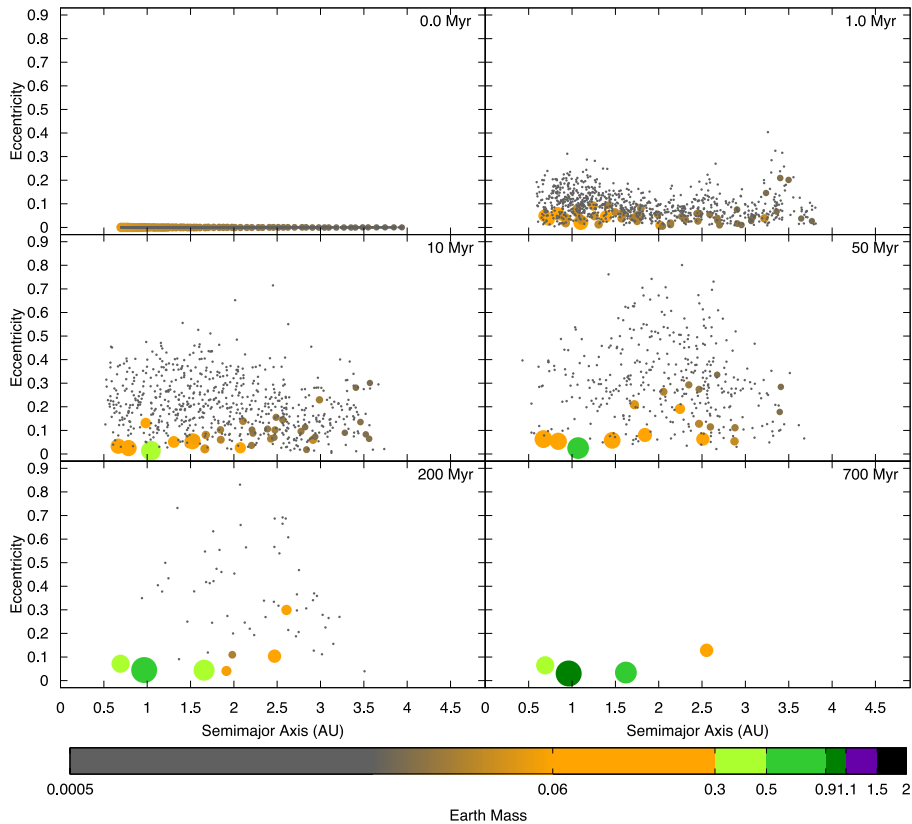


Figure 3. Snapshots of the dynamical evolution of one simulation where $x = 2.5$. Jupiter and Saturn are initially as in the Nice model II. The size of each body corresponds to its relative physical size and is scaled as $M^{1/3}$, where M is the body mass. However, it is not to scale on the x -axis. The colour coding gives the range of mass which each body belongs to.

during the analysis and discussion of our results. The initial orbital inclinations of planetesimals and planetary embryos were chosen randomly from the range of 10^{-4} – 10^{-3} deg, and their other orbital angles were randomized. Their initial eccentricities are set equal to zero.

Our simulations also included fully formed Jupiter and Saturn on orbits consistent with the latest version of the Nice model (Levison et al. 2011). Their initial semimajor axes were 5.4 and 7.3 au (2/3 mean motion resonance), respectively. Their eccentricities (and inclinations) were initially $\sim 10^{-2}$ (deg). For each surface density profile we performed 15 simulations with slightly different randomly generated initial conditions for planetary embryos and planetesimals. Collisions between planetary objects are always treated as inelastic mergers that conserve linear momentum. The simulations were integrated for 700 Myr using the Symba integrator (Duncan, Levison & Lee 1998) and a timestep of 6 d. Planetary objects that reach heliocentric distances equal to 120 au are removed from the system. During our simulations, we neglect gas drag and gas-induced migration of the planetary embryos. Namely, we assume that the initial conditions illustrated in Fig. 2 apply at the disappearance of the gas.

3 RESULTS

Figs 3 and 4 show the evolution of accretion in discs with relatively shallow ($x = 2.5$; Fig. 3) and steep ($x = 5.5$; Fig. 4) surface density profiles. A clear difference between these two cases is the accretion time-scale of the final planets. In the simulation with $x = 2.5$, the accretion time-scale is comparable to that observed in traditional

simulations of terrestrial planet formation (over 100 Myr) but in the simulations with $x = 5.5$ growth is much faster (see Raymond et al. 2005, 2007b; Kokubo et al. 2006). This can be understood simply by the fact that steeper discs have more mass in their inner parts, where the accretion time-scale – which depends on the surface density and the dynamical time – is much shorter. The accretion time-scale of terrestrial planets is in this case a few 10 Myr, consistent with some estimates based on radiative chronometers (Yin et al. 2002; Jacobsen 2005; see Section 3.1.2).

The simulation from Fig. 3 formed a planetary system with three planets inside 2 au. The two inner planets at ~ 0.69 and 0.96 au are reasonable Venus and Earth analogues. However, the third planet formed around 1.5 au is about five times more massive than Mars. The Mars problem is pervasive in all simulations with $x = 2.5$. This set of simulations is also characterized by the absence of surviving planetesimals in the region of the asteroid belt and, in contrast, the survival of a larger Mars-size planetary embryos in this same region.

The simulation from Fig. 4 formed a good Mars analogue at 1.58 au with a mass just 30 per cent larger than Mars. As in the previous simulation, a leftover planetary embryo survived beyond the location of Mars, albeit a much smaller one than in the previous case. Embryos stranded in the asteroid belt could easily be removed during a late instability among the giant planets (i.e. the Nice model). However, an embryo more massive than $\sim 0.05 M_{\oplus}$ carves a gap in the asteroid distribution and this gap survives the giant planet instability (O’Brien et al. 2007; Raymond et al. 2009).

The key result from Fig. 4 is that the surviving asteroids have very low eccentricities (and inclinations, as discussed later). Thus, despite forming good Mars analogues, these simulations cannot be

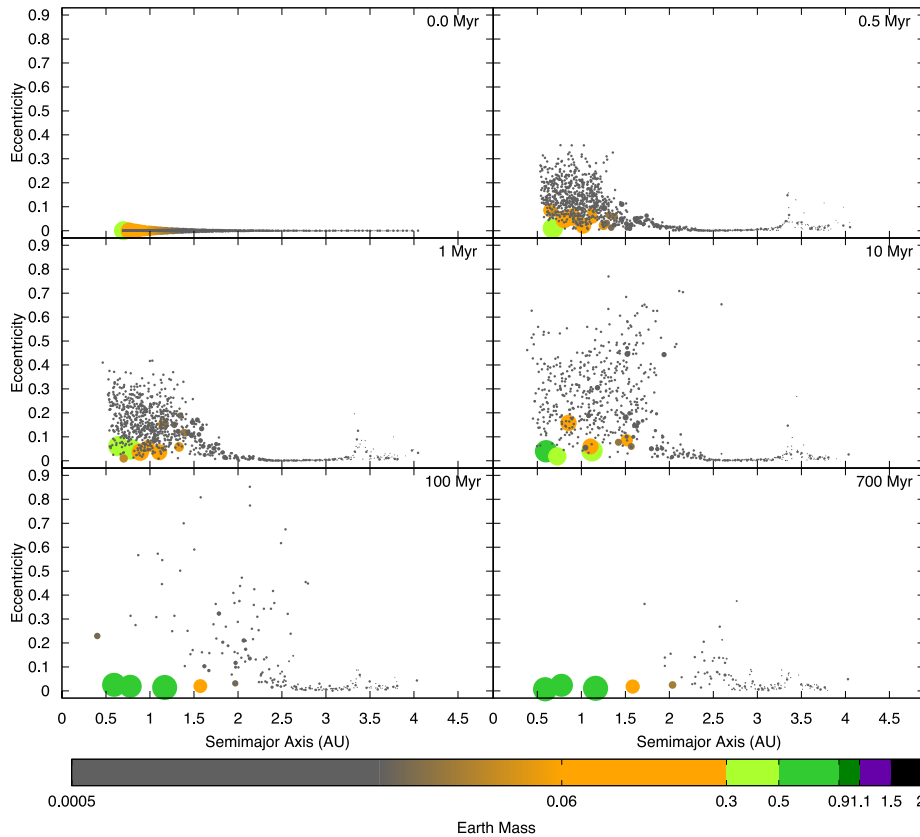


Figure 4. Same as Fig. 3 but for $x = 5.5$.

considered successful. The surviving planetesimals beyond 2.5 au show an orbital eccentricity distribution varying from almost zero to 0.1. However, the observed asteroids have an eccentricity distribution that ranges from 0 to 0.3. The very low asteroidal eccentricities in the simulation are simply because for such a steep surface density profile there is very little mass in the asteroid belt, and this mass is carried only by planetesimals (only about $0.05 M_{\oplus}$). No embryo exists beyond ~ 2.2 au to stir up planetesimals in this simulation. The inner part of the main belt – between 2 and 2.5 au – does have eccentricities up to 0.2, simply because of their closer proximity to planetary embryos. One caveat is that our simulated planetesimals do not self-gravitate and therefore cannot self-excite their eccentricities. We will consider self-interacting planetesimals in the asteroid belt in Section 4.

3.1 The final systems

Figs 5 and 6 show the final orbital and mass distributions of our simulations, sub-divided by their initial disc profile. The total integration time was 700 Myr, roughly corresponding to the start of the late heavy bombardment (LHB; Hartmann et al. 2000; Strom et al. 2005; Chapman, Cohen & Grinspoon 2007; Bottke et al. 2012). We note that our simulations at 400 Myr – the earliest likely start of the LHB (Morbidelli et al. 2012) – are qualitatively the same as at 700 Myr.

3.1.1 The mass and orbital configuration of the planets

It is clear from Fig. 5 that steeper discs (higher x) produce smaller planets around 1.5 au. This is to be expected since steeper discs have

less mass in their outer parts (for a fixed disc mass). Therefore, they are closer to the idealized initial conditions proposed by Hansen (2009), namely a disc truncated at 1 au. The simulations with x of 2.5, 3.5 and 4.5 do not reproduce the terrestrial planets because they produce planets at around 1.5 au that are systematically too massive compared to Mars.² Reasonable Mars analogues only formed in simulations in our steepest disc ($x = 5.5$). This case also produces good Earth and Venus analogues, and 20 per cent of $x = 5.5$ simulations also formed good Mercury analogues. These Mercury analogues are usually leftover planetary embryos that started around 1 au and were gravitationally scattered inwards by growing embryos.

Fig. 6 shows the orbital configuration of the surviving bodies in our simulations. The final orbital distribution of simulated planets roughly matches that of the real terrestrial planets. To perform a quantitative analysis we make use of two useful metrics: the normalized angular momentum deficit (AMD; Laskar 1997) and the radial mass concentration (RMC; Chambers 2001) statistics. The AMD of a planetary system measures the fraction of the planetary system angular momentum missing due to non-circular and non-planar orbits. The AMD is defined as (Laskar 1997)

$$\text{AMD} = \frac{\sum_{j=1}^N [m_j \sqrt{a_j} (1 - \cos i_j \sqrt{1 - e_j^2})]}{\sum_{j=1}^N m_j \sqrt{a_j}}, \quad (1)$$

² In fact, in one of the simulations with $x = 4.5$ we observe the formation of a planet around 1.5 au with mass very similar to Mars (actually slightly smaller). However, in this simulation there is another larger planet around 1.2 au five times larger than Mars.

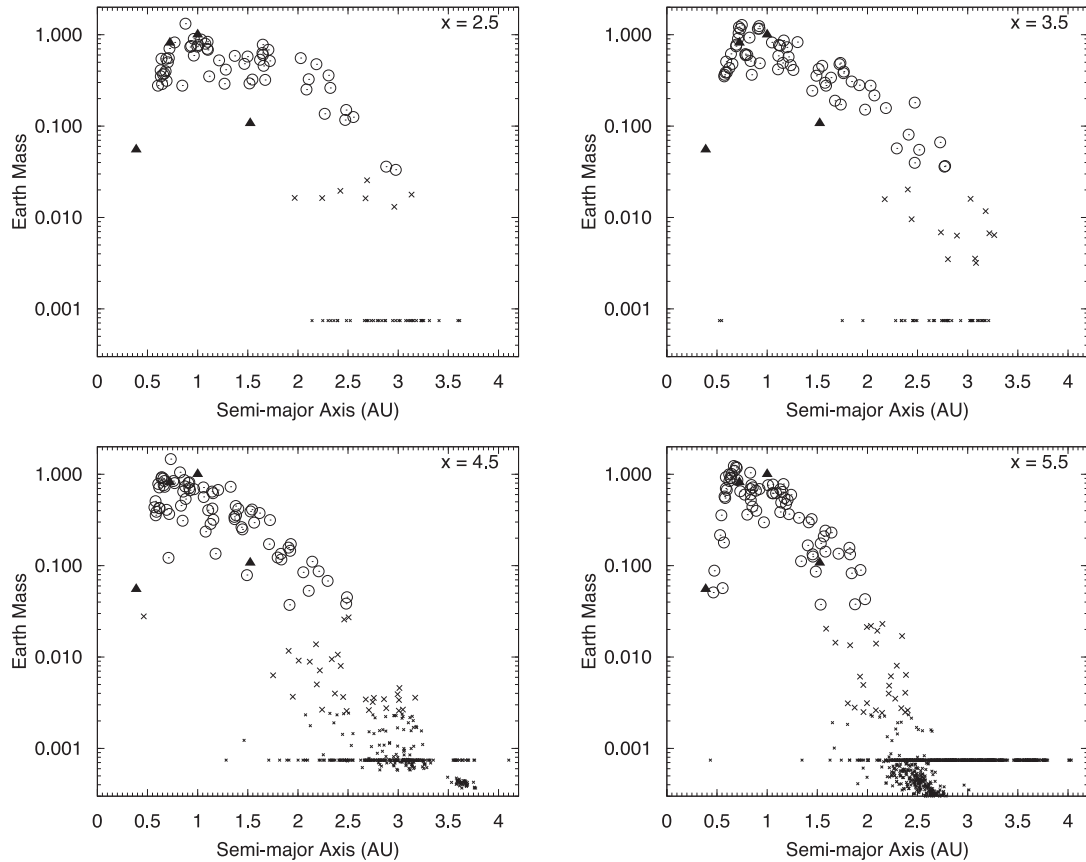


Figure 5. Final distribution of the surviving bodies in our simulations after 700 Myr of integration. Each panel shows the final distribution in a diagram semimajor axis versus mass for all simulations considering the same value of x . The value of x is indicated on the upper-right corner of each panel. Open circles correspond to bodies with masses larger than $0.3 M_{\oplus}$. Smaller bodies are labelled with crosses. The solid triangles represent the inner planets of the Solar system.

where m_j and a_j are the mass and semimajor axis of planet j and N is the number of final bodies. e_j and i_j are the orbital eccentricity and inclination of the planet j .

The RMC of a system measures how the mass is distributed in one region of the system. The value of RMC varies with the semimajor axes of planets (Chambers 1998, 2001; Raymond et al. 2009). It is defined as

$$\text{RMC} = \text{Max} \left(\frac{\sum_{j=1}^N m_j}{\sum_{j=1}^N m_j [\log_{10}(a/a_j)]^2} \right). \quad (2)$$

In these formulas we consider as planets those objects larger than 0.03 Earth masses and orbiting between 0.3 and 2 au.

Table 1 shows the mean number of planets formed in each set of simulations, the mean AMD, RMC and water mass fraction (WMF). This table also provides the range over which these mean values were calculated. Steeper discs produce a larger number of planets per system, in agreement with Raymond et al. (2005). Steeper discs also tend to produce systems that are less dynamically excited. This is a consequence of two effects. First, the initial individual embryo/planetesimal mass ratio in the inner regions of the disc is much higher for steeper discs (this ratio is about 40 [130] at 1 au for $x = 2.5$ [$x = 5.5$]; see Fig. 2), providing stronger dynamical friction, which tends to reduce the planets' final AMD (O'Brien et al. 2006; Raymond et al. 2006). Secondly, shallower discs have more mass in the asteroid belt and a more prolonged last phase of accretion. This

chaotic late phase excites the eccentricities (e.g. Raymond et al. 2014) of the surviving planets while reducing their number.

The final AMD of the simulated planetary systems is also a powerful diagnostic on the adequacy of the chosen initial mass fraction in planetesimals and individual mass ratio between planetary embryos and planetesimals in our simulations (effects of dynamical friction). Simulations of terrestrial planet formation have considered a range in the initial mass fraction of planetesimals, from 0 up to 50 per cent (e.g. Chambers 2001; Raymond et al. 2004; O'Brien et al. 2006; Izidoro et al. 2014; Jacobson & Morbidelli 2014). Given the moderate to low values of AMD of the protoplanetary systems in our simulations, even for the shallower discs, they suggest that our chosen initial mass fraction of planetesimals equal to 30–40 per cent is suitable.

Perhaps the most interesting and surprising result shown in Table 1 is that even our simulations considering $x = 5.5$ were not able to reproduce the large RMC of the Solar system terrestrial planets. The reason for this is that, although very steep discs have by definition a large radial concentration of mass in the inner regions, they tend to produce a large number of planets between 0.3 and 2 au. In fact, most of our simulations produce at least one extra planet between Mars and the inner edge of the asteroid belt (~ 2 au). Simulation with $x = 5.5$, for example, produce on average 4.73 planets per simulation (of course, this is only possible because we have also a very low AMD for these systems which means that planets can stay very close to each other). These 'extra' objects beyond Mars contribute to the observed low values of RMC in these systems.

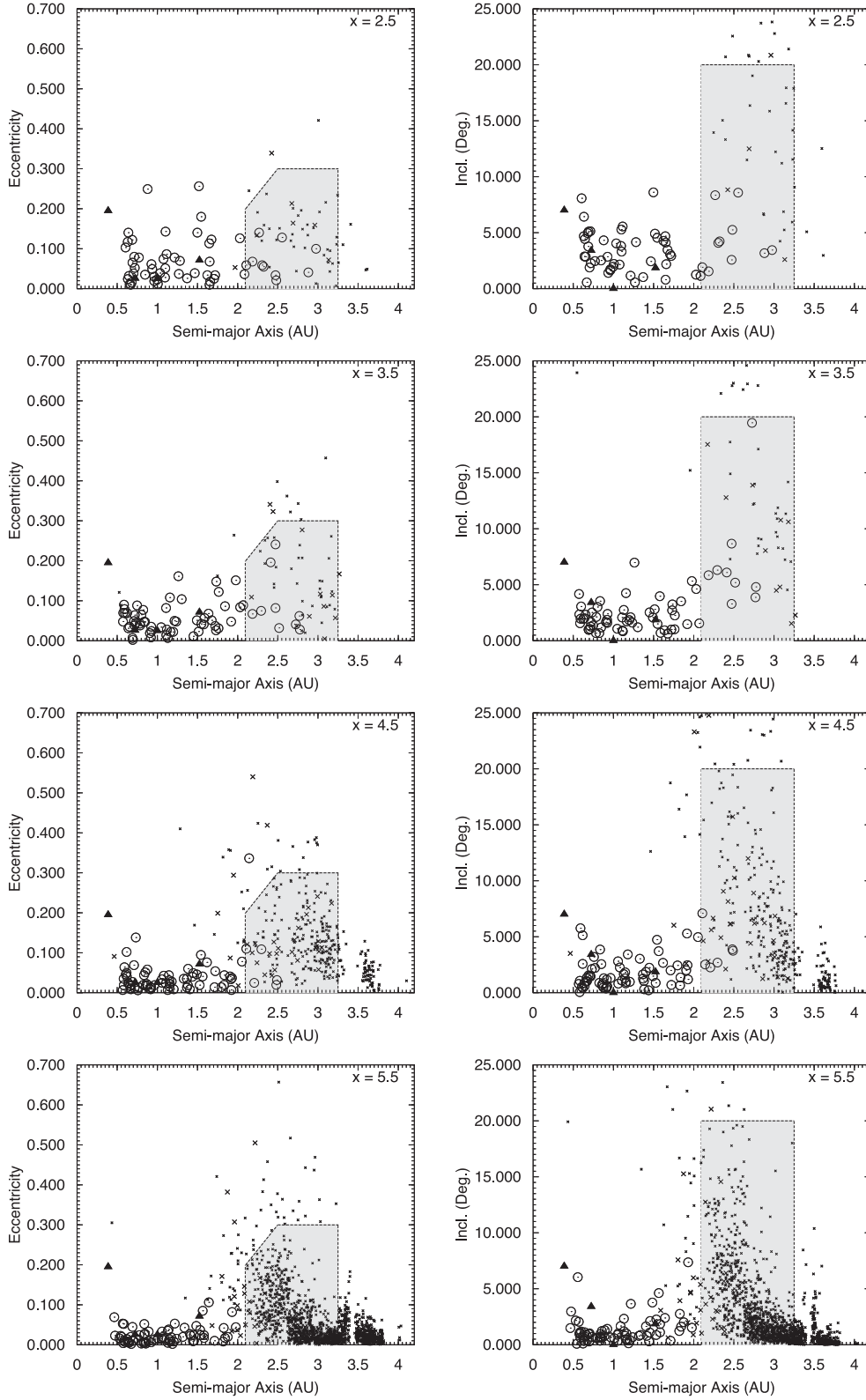


Figure 6. Orbital distributions of the surviving bodies in our simulations after 700 Myr of integration. Open circles correspond to bodies with masses larger than $0.3 M_{\oplus}$. Smaller bodies are labelled with crosses. The left-hand column shows the orbital eccentricity versus semimajor axis while the right-hand column shows orbital inclination versus semimajor axis. The value of x is indicated on the upper-right corner of each panel. The filled triangles represent the inner planets of the Solar system. Orbital inclinations are shown relative to a fiducial plane of reference.

Table 1. Statistical analysis of the results of our simulations and comparison with the Solar system. From left to right the columns are, the slope of the surface density profile, the mean number of planets, the mean AMD, mean RMC and mean WMF, respectively.

x	Mean N	Mean AMD	Mean RMC	Mean WMF
2.5	2.93 (2–4)	0.0047 (0.0003–0.0246)	60.25 (33.33–118.19)	2.64×10^{-3} (9.17×10^{-4} – 4.6×10^{-3})
3.5	3.53 (2–5)	0.0020 (0.0005–0.0058)	50.40 (41.63–62.33)	9.60×10^{-4} (6.46×10^{-4} – 1.84×10^{-3})
4.5	4.33 (2–6)	0.0012 (0.0001–0.0084)	52.23 (42.20–60.47)	3.11×10^{-4} (6.45×10^{-5} – 1.05×10^{-3})
5.5	4.73 (4–6)	0.0004 (0.0001–0.0016)	57.15 (45.87–65.13)	2.81×10^{-5} (2.86×10^{-6} – 9.52×10^{-5})
SS	4	0.0018	89.9	$\sim 10^{-3}$ [Earth ^a]

^aThe amount of water inside the Earth is not very well known. Estimates point to values ranging from 1 to ~ 10 Earth’s ocean (1 Earth’s ocean is 1.4×10^{24} g; see Lécuyer, Gillet & Robert 1998; Marty 2012). A water mass fraction equal 10^{-3} assumes that three oceans are in Earth’s mantle.

However, even if we neglect planets between 1.6 and 2 au when calculating the RMC, the mean value of this quantity in simulations with $x = 5.5$ is 73 and only one planetary system showed an RMC larger than the Solar system value of 89.9.

In fact, no clear trend in RMC is observed for different values of x in Table 1. Moreover, simulations with $x = 2.5$ present a very wide range of RMCs. This may be consequence of stronger dynamical instabilities among forming terrestrial planets which eventually reduce the number of forming planets, sometimes achieving a larger RMC. As we have pointed out before, discs with shallow density profiles are more prone to instabilities because of smaller planetary-embryo/planetesimals individual mass ratio and of the protracted giant impact phase due to the interaction with planetary embryos originally located in the asteroid belt region. As a result, these discs leads to a planetary system with a too large AMD compared to the real Solar system. In summary, we find that no power-law disc profile is able to reproduce simultaneously the AMD and RMC of the real terrestrial planet system. All these results suggest that the high RMC of the terrestrial planets and their low AMD may be a signature of accretion in a very narrow disc with a sharp disc edge truncation necessarily not far from 1 au (Hansen 2009), rather than in a power-law disc, even if very steep. Indeed, in Hansen’s simulations the final systems reproduced the RMC and AMD of the Solar system quite well. The same is true in the Grand Tack simulations (Jacobson & Morbidelli 2014).

3.1.2 Radial mixing and water delivery to Earth

We track the delivery of water from the primordial asteroid belt in to the terrestrial zone (inside 2 au) by inward scattering of planetary embryos and planetesimals. For the initial radial water distribution in the protoplanetary disc we use the model widely used in classical simulations of terrestrial planet formation: planetesimals and protoplanetary embryos inside 2 au are initially dry, those between 2 and 2.5 au carry 0.1 per cent of their total mass in water and, finally, those beyond 2.5 au have initially a WMF equal to 5 per cent (e.g. Raymond et al. 2004; Izidoro et al. 2014). In our simulations, we assume no water loss during impacts or through hydrodynamic escape (see Matsui & Abe 1986 and Marty & Yokochi 2006). The final water content of a planet formed in our simulations is the sum of its initial water content (if any) and the water content of all bodies that hit this respective object during the system evolution. Thus, for our initial water distribution, the WMF of the planets in our simulations probably represents an upper limit (e.g. Genda & Abe 2005).

The mean WMF of the planets formed in our simulations are shown in Table 1. As expected, steeper discs produce drier planets, in agreement with Raymond et al. (2005). In fact, because of the

set-up of radial mass distribution (the total amount of mass is always equal to $\sim 2.5 M_{\oplus}$) of our simulations, steeper discs contain less mass in the outer part of the disc and consequently less water to be delivered to the terrestrial region. To compare quantitatively the results of our simulations with the amount of water in the Earth we assume that the Earth carries three oceans of water in the mantle³ (e.g. Raymond et al. 2009). Under this conservative assumption, only the flattest discs – with $x = 2.5$ and 3.5 – are able to produce, on average, Earth analogues with Earth-like water contents.

It is important to note that time zero of our simulations represent a stage of the formation of terrestrial planets after the gas disc dissipation. The radial mixing and water delivery in our simulations occurs purely due to gravitational interaction among bodies during the evolution of the system. However, recall that the main hypothesis of our work is based on the assumption that the solid mass distribution in the protoplanetary disc is dynamically modified during the gas disc phase by the radial drift of small particles due to gas drag, which can create pile-ups and very steep density profiles. In these circumstances, some level of radial mixing of solid material and also water delivery could have occurred in a very early stage (during the gas disc phase) because of drift of small particles (pebbles or small planetesimals) coming from large distances and entering into the terrestrial region. However, if the snow line had been in the vicinity of 1 au, so to make the disc water-rich near the Earth region, it is difficult to understand why the giant planets formed far out and little mass formed in the asteroid belt (usually the formation of giant planet cores is attributed to snow-line effects). Moreover, it would be strange that asteroids now resident in the inner asteroid belt (S-type, E-type bodies) are predominantly water poor. Our assumption that planetesimals were water-rich only beyond 2.5 au was indeed inspired by the current compositional gradient of asteroids throughout the belt (e.g. DeMeo & Carry 2014).

3.1.3 The last giant impact on Earth

We analyse in this section whether there is any trend or relationship between the timing of the last giant impact on Earth-analogues and the slope of the surface density profile.

Radiative chronometer studies suggest that the last giant collision on Earth happened between 30 and 150 Myr after the formation of the first solids⁴ in the nebula (Yin et al. 2002; Jacobsen 2005;

³ The total amount of water on Earths is debated (Drake & Campins 2006). Estimates suggest a total amount of water in the Earth ranging between 1 and ~ 10 Earth oceans (Lécuyer et al. 1998; Marty 2012).

⁴ The time zero usually is assumed as the time of calcium–aluminium-rich inclusion’s (CAIs) condensation.

Touboul et al. 2007; Allègre, Manhès & G&coumlpel 2008; Taylor, McKeegan & Harrison 2009; Jacobson et al. 2014).

Obviously the time zero of our simulations does not correspond to the time of the CAIs condensation. In our simulations, Jupiter and Saturn are assumed fully formed and the gaseous disc has already dissipated. Thus, it is likely that our scenario represents a stage corresponding to ~ 3 Myr after the time zero (e.g. Raymond et al. 2009). The analysis presented in this section already takes into account this offset in time.

To perform our analysis we define Earth-analogues as planets with masses larger than 0.6 Earth masses, orbiting between 0.7 and 1.2 au and those that suffer the last giant collision between 30 and 150 Myr (e.g. Raymond et al. 2009). We flag as giant collisions those where the scaled impactor mass is larger than 0.026 target masses (or M_{\oplus}).

Fig. 7 shows the timing of last giant impact on all planets larger than 0.6 Earth masses and orbiting between 0.7 and 1.2 au. There, we see a clear trend: planets forming in steeper discs are prone to have the last giant collision earlier. This is consistent with Raymond et al. (2005), who showed that discs with steeper surface density profiles tend to form planets faster. Applying our criterion to flag or not a planet as Earth-analogue, we note that only 40 per cent of the Earth-mass planets formed in our simulations with $x = 5.5$ (which can produce Mars analogues) have the giant impact later than 30 Myr. In many cases, the last giant collision happens much earlier than 30 Myr. Discs with $x = 2.5$, on the other hand, show that about 75 per cent of the forming Earth-mass planets experience their last giant collision later than 150 Myr.

3.1.4 Structure of the asteroid belt

A successful model for terrestrial planet formation must also be consistent with the structure of the asteroid belt. In fact, the asteroid belt provides a rich set of constraints for models of the evolution of the Solar system (Morbidelli et al. 2015). The present-day asteroid belt only contains roughly 10^{-3} Earth masses. Fig. 1 shows the current distribution of asteroids with magnitude smaller than $H < 9.7$. This cut-off corresponds to objects larger than 50 km which safely make this sample free of observational biases (Jedicke, Larsen & Spahr 2002). In addition, such large objects are unaffected for non-gravitational forces and they are unlikely to have been produced in family formation events. As shown in Fig. 1, the main belt basically spans from 2.1 to 3.3 au and is roughly filled by orbital eccentricities from 0 to 0.3 and inclinations from 0 to 20 deg. Thus asteroids are, on average, far more excited than the terrestrial planets. The challenge is to simultaneously explain the belt's low mass and orbital excitation.

We now analyse the formation of the asteroid belt in our N -body simulations. Given the initial conditions that we have adopted for Jupiter and Saturn (different from the current orbits), our simulations implicitly assume that there will be a later instability in the giant planets' orbits. Thus, it is natural to wonder whether such a late instability would alter the final asteroid belt's structure. The version of the Nice model that is consistent with observations is called the jumping-Jupiter scenario (Brasser et al. 2009; Morbidelli et al. 2009b; Nesvorný & Morbidelli 2012), and invokes a scattering event between Jupiter and an ice giant planet (Uranus, Neptune or a rogue planet of comparable mass). In this model the orbital inclinations of the asteroids are only moderately altered (with a root-mean-square change of the inclination of only 4 deg; Morbidelli et al. 2010). The subsequent 3.5 Gyr would have brought even smaller changes in the

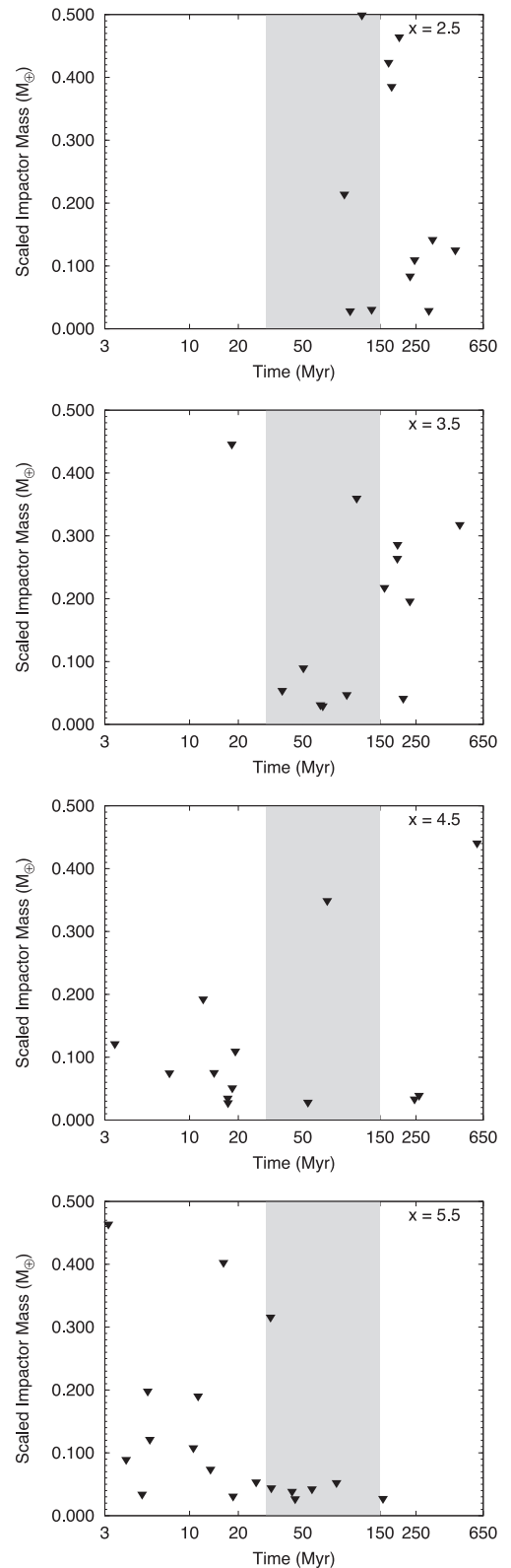


Figure 7. Timing of last giant collision on Earth-analogues for all our simulations. The surface density profile is indicated on the upper-right corner of each plot. The grey area shows the expected range (30–150 Myr) for the last giant impact on Earth derived from cosmochemical studies.

Table 2. Statistical analysis of the orbital architecture of the asteroid belt. Comparison of the rms of the orbital eccentricities and inclinations of the real and simulated populations of asteroids and use of the Anderson–Darling statistical test to compare the different populations. The columns are as follows: root-mean-square eccentricity (e_{rms}), inclination (i_{rms}) and number (N) of planetesimals (or asteroids). Each of these columns are divided in two sub-columns showing the range of semimajor axis of the asteroids taken in each sample for the calculation of e_{rms} , i_{rms} and N . The respective sub-populations are called A which contains only those bodies orbiting with semimajor axis between 2.1 and 2.7 au and sub-population B which contains bodies orbiting with semimajor axis between 2.7 and 3.2 au. In the left, we have the real and simulated population of asteroids (power index x). In the first column from left, we also show the corresponding time, representing a snapshot of the system dynamical state, from which our statistics are computed. Each table entry in parentheses is the p -value of the respective simulated sample when compared with the real population (Fig. 1) using the Anderson–Darling statistical test.

	Region	e_{rms}		i_{rms} (deg)		N	
		2.1–2.7 au [A]	2.7–3.2 au [B]	2.1–2.7 au [A]	2.7–3.2 au [B]	2.1–2.7 au [A]	2.7–3.2 au [B]
Real population		0.18	0.16	11.2	12.6	173	352
power index x /time							
2.5/300 Myr		0.21 (0.07)	0.17 (0.23)	20.1 (0)	16.4 (10^{-4})	33	59
2.5/700 Myr		0.19 (0.68)	0.13 (0.18)	19.3 (10^{-5})	16.5 (0.01)	13	18
3.5/300 Myr		0.21 (0.01)	0.19 (10^{-3})	18.3 (0)	12.1 (0.67)	64	84
3.5/700 Myr		0.22 (0.07)	0.15 (0.17)	19.7 (10^{-5})	12.8 (0.29)	10	28
4.5/300 Myr		0.20 (0.17)	0.15 (10^{-3})	12.3 (0.07)	7.13 (0.0)	189	277
4.5/700 Myr		0.17 (0.22)	0.15 (0.13)	13.4 (10^{-3})	7.8 (0.0)	55	162
5.5/300 Myr		0.12 (0.0)	0.04 (0.0)	6.9 (0.0)	2.2 (0.0)	546	623
5.5/700 Myr		0.13 (0.0)	0.05 (0.0)	7.8 (0.0)	2.2 (0.0)	382	620

asteroid inclination distribution. Thus, in our analysis we compare directly the inclination and eccentricity distribution of the asteroids resulting from our simulations with the observed distribution, with the understanding that, if the simulated inclination distribution is too cold, it is highly unlikely that the subsequent evolution would reconcile the two distributions. To consolidate this claim, in Section 4.3, we perform numerical simulations to illustrate how the jumping-Jupiter scenario would affect the orbital distribution of our final systems. We note that the inverse is not necessarily true: a modestly overexcited asteroid belt could in principle be dynamically ‘calmed’ during the Nice model instability by removal of the most excited bodies (Deienno et al. 2015).

Our main result is that only a subset of our simulations adequately excite the asteroid belt. Fig. 6 shows the final orbital distribution of the protoplanetary bodies in all simulations after 700 Myr. For $x = 2.5, 3.5$ and 4.5 planetary embryos as large as Mars were initially present in the asteroid belt (see Fig. 2), and many of these survive in the main belt (see also Fig. 5). These planetary embryos interact gravitationally with planetesimals increasing planetesimals’ velocity dispersions and generating higher eccentricities or/and inclinations. Of course, it is inconsistent with observations for large embryos to survive in the main belt because they would produce observable gaps in the resulting asteroid distribution (O’Brien et al. 2006; Raymond et al. 2009). This indeed happens for $x = 2.5, 3.5$ and 4.5 .

For $x = 2.5$, most of the bodies that survived between 2 and 3 au are large planetary embryos. Planetesimals that survived in this region have very high orbital inclinations and orbital eccentricities above 0.1. Only the very outer part of the main belt contains a significant planetesimal population. The situation is similar for simulations with $x = 3.5$. These cases provide a reasonable match to the orbital distribution of bodies contained within the belt but the existence of planetary embryos makes them inconsistent with today’s asteroid belt.

In the simulations with $x = 4.5$ many planetary embryos survived interior to 2.5 au. The region exterior to 2.5 au is mostly populated by planetesimals (and some <Moon-mass planetary embryos; see Fig. 5) with a broad range of excitations, although there is a modest deficit of high-inclination planetesimals in the outer main belt. This set of simulations provides the closest match to observations. Re-

member, however, that in addition to the remaining long-lived large planetary embryos frequently observed in the asteroid belt (Figs 5 and 6), in this disc planets around 1.5 au are systematically about five times more massive than Mars (Fig. 5).

The simulations with $x = 5.5$ (the only ones that reproduce the small mass of Mars) suffer from a severe underexcitation of the asteroid belt (Fig. 6). Planetesimals initially between 2.1 and 2.5 au gain eccentricities up to 0.2–0.3 and inclinations up to 15 deg. However, planetesimals beyond 2.5 au are characterized by $e < 0.1$ and $i < 10^\circ$. Indeed, the typical inclination in the outer belt is less than 3° . This is certainly inconsistent with the real population of asteroids. We stress that any initial disc mass distribution which is not a power law but predicts less mass in the asteroid belt than our disc with surface density profile proportional to $r^{-5.5}$ (e.g. Hansen 2009) would obviously have the same problem.

We now quantitatively compare the results of our simulations with the real asteroid belt. Table 2 shows the root mean square (rms) of the orbital eccentricity (e_{rms}) and inclination (i_{rms}) of real and simulated populations.

To perform our analysis we divided the real main belt in two sub-populations and calculated the quantities (e_{rms} and i_{rms}) in each. Bodies belonging to the sub-population A have semimajor axis between 2.1 and 2.7 au and those from the sub-population B have semimajor axis ranging from 2.7 to 3.2 au. We divide the simulated belt populations following these same criteria. We only included simulated planetesimals in the asteroid belt region for which the orbital eccentricity is smaller than 0.35, the orbital inclination is smaller than 28 deg (see Fig. 1) and which are confined within either region A or B. In principle, these cut-off values were chosen using as inspiration the ‘edges’ of the real population of asteroids in the main belt with magnitude $H < 9.7$ (see Fig. 1). Planetesimals with e or i (significantly) higher than those of the real population are eventually removed from the system due to close encounters with the giant planets or with the forming terrestrial planets. This is observed in our simulations. For example, we computed the fraction of planetesimals between 2.1 and 2.7 au that have orbital eccentricity and inclination higher than our cut-off values, at two different snapshots in time, at 200 and 700 Myr. Our results show that, at 200 Myr, 40 and 27 per cent of planetesimals between 2.1 and 2.7 au have eccentricities larger than 0.35 in simulations with $x = 2.5$

and 3.5, respectively. At 700 Myr, however, for discs with $x = 2.5$ this number is reduced to 0 per cent (meaning that there is no planetesimal with orbital eccentricity higher than our cut-off value – they have been mostly ejected from the system or collided with the star). For $x = 3.5$ this number drops to 20 per cent at 700 Myr. Thus, ignoring such bodies in our analysis is adequate since we avoid biasing our conclusions in this sense.

Table 2 also shows the size of our samples, i.e. the sizes of the populations of real and simulated asteroids composing the regions A and B in our analysis. For robustness, our analysis of the (e_{rms}) and (i_{rms}) is presented together with a statistical test. We use the Anderson–Darling (AD) statistical test, a more sensitive and powerful refinement of the popular Kolmogorov–Smirnov statistical test (Feigelson & Jogesh Babu 2012). The AD test is often used to decide whether two random samples have the same statistical distribution. A low probability (or p -value) means that the samples are dissimilar and unit probability means they are the same. We apply the AD test separately to sub-populations A and B and use a significance level to discard the similarity among samples smaller than 3 per cent.

Table 2 summarizes our analysis. Before comparing the simulated and real belt populations let us compare regions A and B of the real belt population. The e_{rms} of asteroids in regions A and B of the real main belt are similar. The same trend is observed in i_{rms} values. The excitation level over the entire asteroid belt is equally balanced, providing a strong constraint for our simulations. Because the final number of objects in regions A and B are relatively small for those simulations with $x = 2.5$ and 3.5 (after 700 Myr of integration) we also apply our statistical tests using the system dynamical states at 300 Myr. Around 300 Myr, the number of planetesimals in the asteroid belt region is higher than that at 700 Myr but most of the planets are almost or fully formed (e.g. Fig. 3) which supports this choice. This secondary analysis is also certainly important in the sense of certifying that our statistical analyses are not biased by the smaller number of surviving objects in these systems.

The level of eccentricity excitation of the planetesimals in the A region, at both 300 and 700 Myr, are comparable for simulations with $x = 2.5$, 3.5 and 4.5. This result is counter-intuitive since it does not reflect the fact that shallow discs have more mass and larger planetary embryos in the asteroid belt and so should dynamically excite planetesimals more strongly than steeper discs (Figs 5 and 6). We do indeed find that simulations with shallower disc profiles have a larger proportion of planetesimals that are overexcited with respect to our asteroid belt e and i cut-offs, and a faster decay in this population. However, the rms excitation of the planetesimals within the defined belt remains (for some discs; see Table 2) similar. We suspect that this may be linked with a combination of small number statistics and dynamical excitation by remnant planetary embryos in the belt.

The AD test applied to the A region of our simulations with $x = 2.5$, 3.5 and 4.5 does not reject the hypothesis that these populations have the same statistical distribution than the real belt. In fact the p -value in these cases are about 68, 7 and 22 per cent for simulations with $x = 2.5$, 3.5 and 4.5, respectively. However, the disc with $x = 5.5$ is dynamically much colder than the real population. The AD test rejects with great confidence the possibility that this simulated population of asteroids match the real one.

As may be expected, there are some modest differences for the values of e_{rms} and i_{rms} (and consequently the p -values) shown in Table 2 at 300 and 700 Myr. It is natural to expect that a system may be dynamically excited over time as far there is a mechanism to do so. However, we also note that a system may be dynamically ‘calmed’ if a fraction of the objects with (very) eccentric/inclined

orbits are removed from the system while those with dynamically colder orbits tend to be kept in the system. This is observed, for example, for our simulations with $x = 2.5$ (the e_{rms} of the A and B regions are smaller at 700 Myr than at 300 Myr). However, to summarize, none of our discs had its dynamical state altered from 300 to 700 Myr up to the point that the simulated asteroid belt matches the real one.

Our simulated population of outer main belt (region B) asteroids in discs with $x = 2.5$, 3.5 and 4.5 have eccentricity distributions that match the real outer belt. However, the B population of our discs with $x = 5.5$, are statistically underexcited compared with the real belt.

The real asteroid belt’s inclination distribution turns out to be the most difficult constraint for the simulations, in particular the similar i_{rms} values of the inner and outer main belt. Our simulations with $x = 2.5$, 3.5 and 4.5 produce inner main belts (region A) with inclinations that are statistically overexcited compared with the real one. The simulation with $x = 5.5$, on the other hand, drastically underexcites the inclinations of the inner main belt. In the outer main belt (region B), only the simulations with $x = 3.5$ provide a reasonable match to the actual asteroid inclinations. Simulations with steeper surface density profiles ($x = 4.5$ and 5.5) underexcite the outer main belt’s inclinations while simulations with shallower profiles ($x = 2.5$) overexcite the inclinations.

To summarize, none of our simulations produced an asteroid belt with the same level of dynamical excitation as the present-day main belt. Of course, we need to keep in mind that the end state of our simulations does not correspond to the present day. Rather, it only corresponds to the start of the late instability in the giant planets’ orbits envisioned by the Nice model, which certainly acted to modestly deplete and reshuffle the asteroid belts’ orbits (Morbidelli et al. 2010; Deienno et al. 2015). In addition, there are other potential mechanisms of excitation that we have not yet considered.

4 MECHANISMS OF DYNAMICAL EXCITATION OF THE ASTEROID BELT

As shown before, the only disc capable of producing a Mars-mass planet around 1.5 au results in an asteroid belt with an orbital distribution way too cold compared to the real one. Complementing the analysis disclosed previously, we present in Section 4.1 the results of numerical simulations including N -body self-interacting planetesimals and results of semi-analytical calculations describing the gravitational stirring among a large number of planetesimals.

4.1 Self-interacting planetesimals in the asteroid belt

The key open question at this point in our study is whether, in the absence of planetary embryos, the asteroid belt can self-excite. This is the scenario represented by our simulations with $x = 5.5$. In those simulations the region beyond 2.5 au contains many planetesimals but they have very low eccentricities and inclinations. However, the simulations did not include interactions between planetesimals, which should act a source of viscous stirring to increase the orbital excitation. We tackled this problem using both an N -body and a particle-in-a-box approach.

4.1.1 N -body numerical simulations

In this scenario, we assume that the region of the disc beyond 2 au is populated only by planetesimals. We performed a simulation

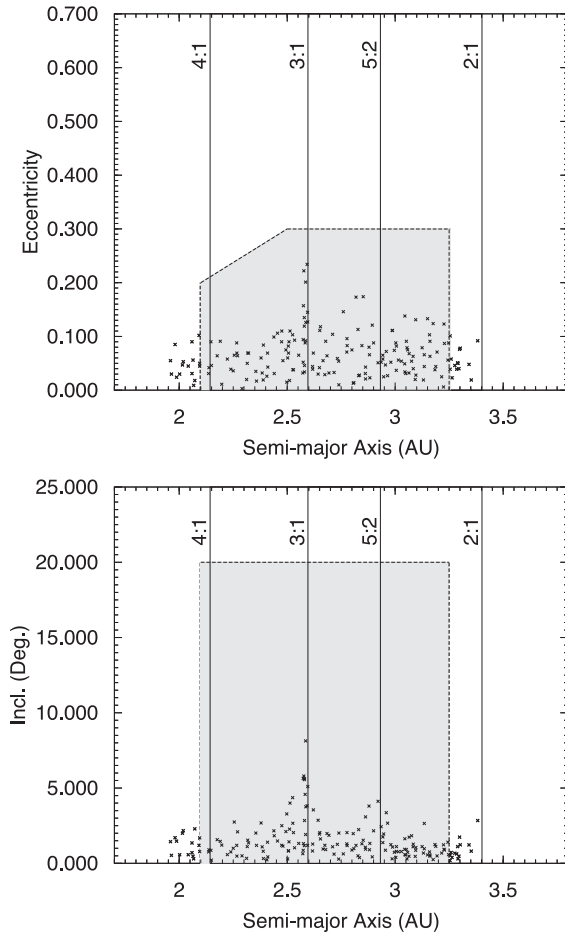


Figure 8. Final state of one simulation considering self-interacting planetesimals distributed between 2 and 3.5 au after 200 Myr of integration. Each planetesimal has an initial small mass equal to $1.3 \times 10^{-4} M_{\oplus}$. The upper plot shows eccentricity versus semimajor axis. The lower plot shows orbital inclination versus semimajor axis. The vertical lines show the locations of mean motion resonances with Jupiter.

containing 300 asteroids with individual masses of $1.3 \times 10^{-4} M_{\oplus}$. This translates to objects with diameters of about ~ 1000 km. The total mass is $0.04 M_{\oplus}$, which is about 80 times the current belt mass. We uniformly distributed these objects between 2 and 3.5 au with initially circular orbits and inclinations chosen randomly from the range of 10^{-4} – 10^{-3} deg (mean anomalies and longitudes of ascending nodes were randomized). Jupiter and Saturn have the same orbital configuration as our previous simulations. We have numerically integrated the dynamical evolution of these bodies for 200 Myr.

Fig. 8 shows the self-excitation of the simulated asteroid belt after 200 Myr. Orbital eccentricities of asteroids are excited up to about 0.1 and orbital inclination to no more than 2° – 4° . And these values of eccentricity and inclination should be considered as upper limits, because the self-excitation does not lead to a substantial loss of bodies, as it is the case here, and the belt could not contain so much mass in large asteroids. Thus, the real asteroid self-excitation would have been even smaller.

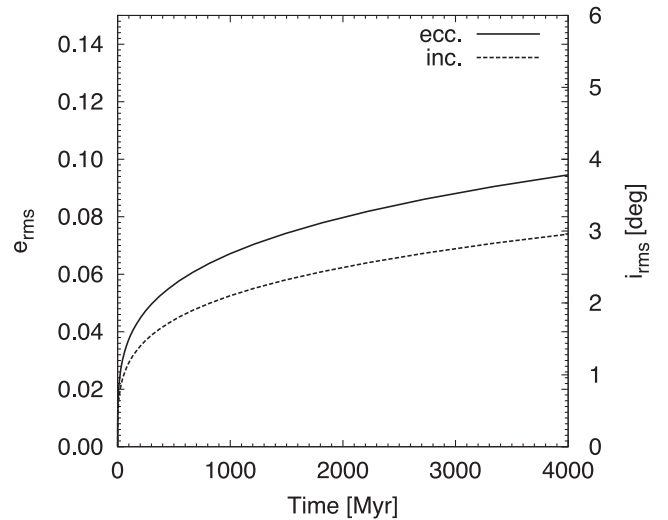


Figure 9. Evolution of rms eccentricity and inclination of a population of 1000 Ceres-mass planetesimals for 4 Gyr obtained from statistical calculations.

4.1.2 Semi-analytical calculations: Particles in a box approximation

We also performed an analysis using the particle-in-a-box approximation (e.g. Safronov 1969; Greenberg et al. 1978; Wetherill & Stewart 1989, 1993; Weidenschilling et al. 1997; Kenyon & Luu 1999; Stewart & Ida 2000; Morbidelli et al. 2009a). We used a very similar set-up to the N -body simulations. Our calculation describes the gravitational stirring among a large number of planetesimals in the region of the asteroid belt, using the Boulder code (Morbidelli et al. 2009a). As in the previous case the individual masses of planetesimals are purposely chosen to represent an extreme case compared to the expected size frequency distribution for the primordial population of asteroids in the region (Bottke et al. 2005; Morbidelli et al. 2009a; Weidenschilling 2011). Our dynamical system is composed by a one-component population of 1000 Ceres-mass planetesimals with diameter of about 1000 km distributed from 2 to 3.2 au. The orbital eccentricities and inclinations of these bodies are set initially to be equal to $e_0 = 0.0001$ and $i_0 \simeq 2.8 \times 10^{-3}$ deg. We follow the evolution of the rms eccentricity and inclination of this population of objects for 4 Gyr, nearly the age of the Solar system.

Fig. 9 shows the outcome of this calculation. As expected, the rms eccentricity and inclination of this population is much smaller than that of the real population of the bodies in the asteroid belt. The e_{rms} of this populations stays below 0.1 while the orbital inclination is smaller than 3° . This is a robust result given that the amount of mass carried out by 1000 Ceres mass is about ~ 0.2 Earth masses. However, the formation of Mars analogues around 1.5 au is only possible in those discs as steep as $r^{-5.5}$. But the amount of mass between 2 and 3.2 au in these discs (see Section 2) is only about ~ 0.03 Earth masses. Thus, the results from Fig. 9 are firm upper limits on the disc's self-excitation.

It is important also to stress that in our semi-analytical calculation we neglect mass accretion, fragmentation and erosion due to collisions between planetesimals. If taken into account, these effects in general should only slightly change the evolution of e_{rms} and i_{rms} of this population (e.g. Stewart & Ida 2000). However, their effects could become important in cases where there is mass growth and formation of very large bodies during the time-scale of the

evolution of the system. In this case, our calculation would probably underestimate e_{rms} and i_{rms} since we are using a one-component population of planetesimals (during all the 4 Gyr the planetesimals have the same mass). In our application, however, this seems to be an adequate approach since we do not see in the asteroid belt today anybody larger than Ceres.

Of course, our statistical calculation does not take into account the gravitational perturbation from the giant planets. However, because they have initially almost coplanar and circular orbits their secular effects on the planetesimals are very small (see Raymond et al. 2009 and our Fig. 8). In this case, only those bodies near mean motion resonances should gain modest inclinations and eccentricities.

4.2 Late dynamical instabilities and migration of giant planets

The giant planet configuration assumed in our simulations represents an epoch prior to the late stage of dynamical instabilities of the giant planets. In our simulations, only the very steepest power-law disc (with $x = 5.5$) formed viable Mars analogues, but the asteroid belts in these simulations were underexcited compared with the present-day belt. But could the later stage of dynamical instability among the giant planets solve this issue? To address this question we performed simulations to analyse how the final state of the planetary systems produced in discs with $x = 5.5$ is affected after the migration of Jupiter and Saturn within the ‘jumping-Jupiter’ scenario (Brasser et al. 2009; Morbidelli et al. 2009b; Nesvorný & Morbidelli 2012). Specially, we focus on the evolution of the orbital inclination of the asteroid belt.

To mimic the evolution of Jupiter and Saturn in the jumping-Jupiter scenario, for simplicity, we assume that Jupiter and Saturn instantaneously evolve from the final state of our simulations of terrestrial planet formation, where they still have almost coplanar and circular resonant orbits (at $t = 700$ Myr) to their current orbits. We use this approach for two reasons. First, it is computationally cheap. Secondly, this assumption avoids many difficulties that would emanate from using a forced migration for these planets. A poorly controlled forced migration for Jupiter and Saturn could bias our analysis and conclusions. In contrast, this approach allows us to assess the essential signatures printed in the asteroid belt by the jumping Jupiter. Simulations of the jumping-Jupiter process (Nesvorný & Morbidelli 2012) show that it is unlikely that Jupiter passed through a dynamical phase with an inclination significantly larger than its current one. Thus, substituting the original Jovian orbit with quasi-null inclination with the current orbit, as we do here, is a good approximation for the jumping-Jupiter evolution, at least as far as the inclination excitation is concerned.

Fig. 10 shows the final state of all our simulations, with $x = 5.5$, 30 Myr after the jumping-Jupiter-like evolution of Jupiter and Saturn. We note that the final orbital distribution of bodies after the jumping-Jupiter-like evolution of Jupiter and Saturn does not change qualitatively. In fact, the strongest effect appears in the orbital eccentricity distribution, but still planetesimals beyond 2.5–2.8 au have eccentricities ~ 0.1 or smaller. The orbital inclinations of bodies beyond 2.5–2.8 au also are only weakly affected.

We stress that even if Jupiter and Saturn had reached their current orbits in a much smoother and slow migration fashion (Tsiganis et al. 2005; Minton & Malhotra 2009, 2011) than predicted by the jumping-Jupiter scenario, the orbital inclination distribution of simulated bodies beyond ~ 2.8 au (simulations with $x = 5.5$) would remain very dynamically cold. This is because planetesimals beyond 2.8 au would preserve their initial inclination distribution because they are not greatly affected by the sweeping of secular res-

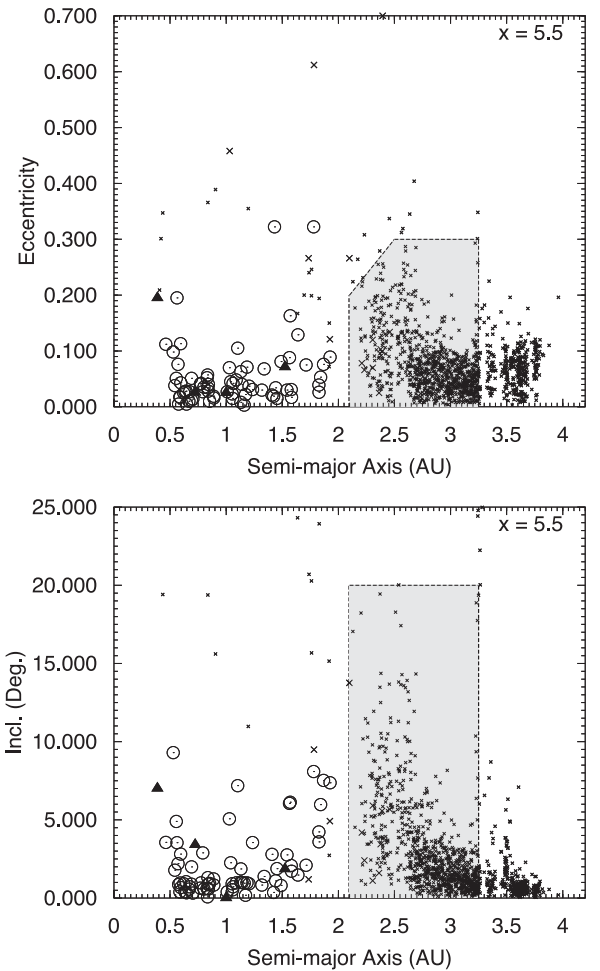


Figure 10. Same as that for Fig. 6 but after the jumping-Jupiter-like evolution of Jupiter and Saturn.

onances during the giant planet migration (e.g. fig. 4 in Morbidelli et al. 2010). In contrast, most of the planetesimals/asteroids inside 2.8 au would have their orbital inclinations increased significantly. Slow migration of giant planets also comes with the drawback of implying a slow sweeping of secular resonances across the asteroid belt and terrestrial region which likely make the terrestrial planets too dynamically excited in the end (Brasser et al. 2009; Agnor & Lin 2012).

4.3 Effects of the initial orbits of the giant planets.

In this paper we have assumed that Jupiter and Saturn were on nearly circular and coplanar orbits during the late stages of terrestrial accretion. Yet the giant planets’ orbits at early times are poorly constrained. There are two qualitatively different schools of thought (see recent reviews by Morbidelli et al. 2012, Raymond et al. 2014 and Raymond & Morbidelli 2014 for more detailed discussions).

The first model proposes that, when the gaseous protoplanetary disc dissipated, the giant planets were stranded on orbits sculpted by planet–disc interactions. The most likely orbital configuration is a low-eccentricity, low-inclination 3:2 mean motion resonance between Jupiter and Saturn (Masset & Snellgrove 2001; Morbidelli & Crida 2007; Pierens & Nelson 2008; although a 2:1 resonance is also possible – Pierens et al. 2014). The giant planets acquired their current orbits during a late phase of orbital instability, i.e. the Nice

model (Morbidelli et al. 2007; Batygin & Brown 2010; Levison et al. 2011; Nesvorný & Morbidelli 2012). Within a standard, $\Sigma \propto r^{-1.5}$ disc, this model systematically produces unrealistically large Mars analogues (Chambers 2001; Raymond et al. 2009; Morishima et al. 2010). This problem motivated the development of the Grand Tack scenario (Walsh et al. 2011).

The second model proposes that Jupiter and Saturn were on or close to their current orbits from very early times (we refer the reader to Raymond et al. 2009 for a discussion), prior to the terrestrial planet accretion phase (e.g. Chambers 2001). This could be the case if the giant planet instability occurred as soon as the gas was removed from the protoplanetary disc, rather than at a late time. In this case, obviously, the giant planet instability cannot be the source of the LHB of the terrestrial planets, unlike postulated in the Nice model (Gomes et al. 2005; Levison et al. 2011). However, the very existence and the nature of the LHB is still debated (e.g. Hartmann et al. 2000), so this possibility cannot be excluded with absolute confidence.

In practice, though, the giant planets' eccentricities were damped by scattering planetary embryos and planetesimals during terrestrial accretion. So in order to end up on their current orbits the giant planets' early orbits must have been *more* eccentric than their current ones, with $e_{J,S} \approx 0.07\text{--}0.10$; this configuration was called EEJS in Raymond et al. (2009) and Morishima et al. (2010). The problem is that no simulations of the giant planet instability (either in its 'early' or 'late' versions) that successfully reproduced the outer Solar system ever succeeded in producing an orbit of Jupiter about twice more excited in eccentricity than the current one (Nesvorný 2011; Nesvorný & Morbidelli 2012). For this (and other) reasons, the EEJS configuration is not considered realistic in Raymond et al. (2009).

If one neglects these issues, the EEJS model can naturally produce a small Mars starting from a shallow surface density profile of the solid distribution ($x \sim 1.5$), because material is removed in the vicinity of the Mars region by the action of the ν_6 secular resonance at 2.1 au, which is initially super-strong in view of the larger eccentricity of Jupiter (Raymond et al. 2009). In this scenario, a lot of mass beyond ~ 2.5 au is also removed from the system because of the strong gravitational perturbation of Jupiter and Saturn.

However, for the goal of the current paper, which is to investigate whether a steep radial distribution of solids could explain simultaneously the small mass of Mars and the asteroid belt orbital properties, the EEJS set-up would not help. In fact, a steep surface density profile would have a small mass in the Mars region to start with, and the strong local depletion acted by the ν_6 resonance would further reduce the mass and in the end Mars would be too small. In terms of the asteroid belt, a more eccentric Jupiter would not substantially enhance the dynamical excitation of the asteroid belt, away from the ν_6 resonance (Fig. 11).

4.4 An alternate model with a localized mass depletion (Izidoro et al. 2014)

This paper was strongly motivated by the results of Izidoro et al. (2014), who modelled the formation of the terrestrial planets by assuming a local mass depletion between $\sim 1.1\text{--}1.3$ and $\sim 2\text{--}2.1$ au in the original solid mass distribution (Jin et al. 2008). In this section we revisit this scenario discussing the success and limitations of that model.

Izidoro et al. (2014) claimed success in producing Mars-analogues and an excited and depleted asteroid belt in simulations

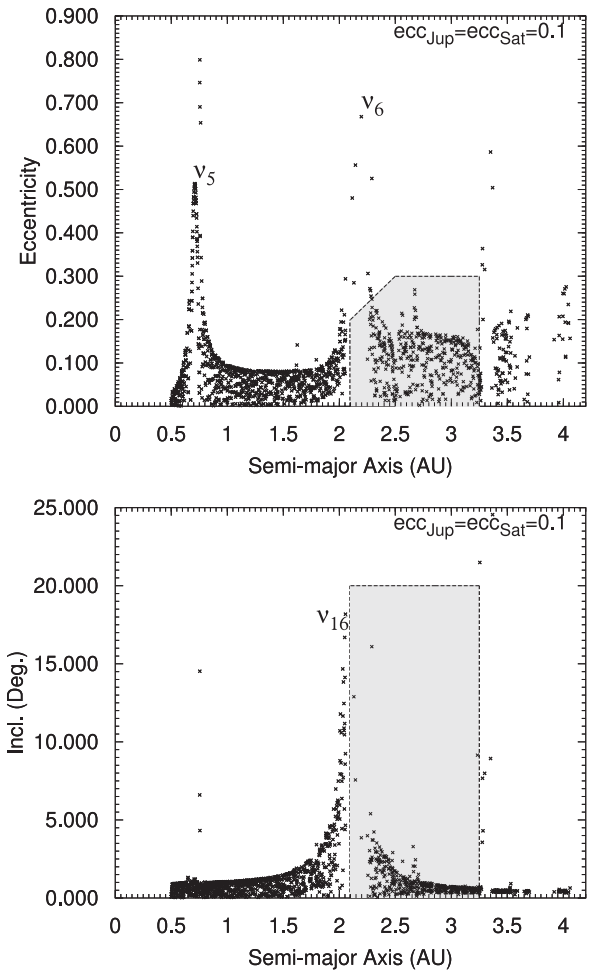


Figure 11. Snapshots at 10 Myr of a simulation considering 2000 test particles distributed between 0.5 and 4 au. Jupiter and Saturn are initially as in their current orbits but with slightly higher orbital eccentricities. Secular resonance locations are also shown.

considering initially Jupiter and Saturn as in their current orbits and a high mass depletion between 1.3 and 2 au. One of the drawback in the results of Izidoro et al. (2014), however, is that simulations considering the giant planets in almost circular and coplanar orbits – as envisioned by models of the dynamical evolution of the Solar system – failed to produce Mars-mass planets around 1.5 au (Morbidelli et al. 2012; Raymond et al. 2014; Raymond & Morbidelli 2014). In Izidoro et al. (2014) the eccentric orbits of the giant planets, combined with an assumed severe initial local mass-depletion at around 1.5 au (between 1.3 and 2 au), are the key to produce Mars-analogues. This is because the giant planets' gravitational perturbations are stronger in this case (relative to the case where Jupiter and Saturn have almost circular and coplanar orbits) and help quickly removing a significant fraction of the ~ 2 Earth masses of solid material initially beyond 2 au. This avoids that planets (Mars-analogues candidates) forming around 1.5 au accrete too much mass. However, Jupiter and Saturn most likely had different orbits at the time the terrestrial planets were still forming (Section 4.3).

Similarly to the results of our shallow discs and other classical simulations (Chambers 2001; O'Brien et al. 2006; Raymond et al. 2006), the Izidoro et al. (2014) simulations also produced an

excited and depleted asteroid belt. None the less, qualitatively, the final distribution of bodies in this region only poorly matches the observed one (Izidoro et al. 2014). Another problem of the Izidoro et al. (2014) scenario is that it needed a localized mass depletion in solids around 1.5 au probably too extreme compared to results of disc hydrodynamical simulations (see Jin et al. 2008). The scenario presented in this paper is essentially an updated version of that in Izidoro et al. (2014) using a self-consistent approach which invokes the radial drift and pile-up of mass due to gas drag, leaving a depleted region beyond 1 au. Our current approach is also consistent with an envisioned global picture of the Solar system evolution (Levison et al. 2011), whereas the Izidoro et al. (2014) approach is not.

5 DISCUSSION

The goal of this paper is to evaluate whether the inner Solar system could be reproduced by a disc of terrestrial building blocks with a power-law radial density profile. We have assumed that planetary embryos and planetesimals are distributed from 0.7 to 4 au and have tested four different surface density profiles: $\Sigma \propto r^{-x}$, where $x = 2.5, 3.5, 4.5$ and 5.5 , with a total of $\sim 2.5 M_{\oplus}$ in solids. We adopt the most recent version of the Nice model (Levison et al. 2011), which sets the giant planets' orbits to be resonant, nearly circular and coplanar.

Our main result is that this scenario fails to match the properties of the inner Solar system. A local mass deficit is needed to form a good Mars analogue (e.g. Hansen 2009). Yet a significant mass in planetary embryos in the asteroid belt is required in order to excite the asteroids' eccentricities and, more importantly, inclinations to their current values (e.g. Petit et al. 2001). As expected from these arguments, only our steepest disc profiles ($x = 5.5$) produced good Mars analogues, but those simulations yielded an under-excited asteroid belt. Simulations with flatter disc profiles formed Mars analogues far more massive than the actual planet. Those simulations excited the asteroid belt to roughly the right amount but failed to adequately match observations because too many planetary embryos were stranded in the belt.

Of course, our simulations are limited. For instance, we did not include the effect of non-accretionary collisions (see Leinhardt & Stewart 2012; Stewart & Leinhardt 2012). But recent simulations by Chambers (2013) and Kokubo & Genda (2010) have shown that the effects of fragmentation are minor and we do not expect this to qualitatively affect our results, i.e., to provide a solution to the small Mars problem. We also have considered a single initial total mass for the disc. However, more massive discs only tend to form more massive planets (e.g. Kokubo et al. 2006; Raymond et al. 2007b). While higher mass disc may potentially alleviate at some level the cold asteroid belt issue produced in simulations with steeper discs, a larger initial total amount of mass carried by protoplanetary embryos and planetesimals than we have considered would, in contrast, likely make it more difficult to form Mars-analogues around 1.5 au.

The mass pile-up in the inner Solar system due to gas drag would potentially produce a very steep radial density profile, thus leading to a small Mars and a low-mass asteroid belt. But our simulations imply that the asteroid belt would be too dynamically cold. This mechanism may be valid to explain the *in situ* formation of the cold Kuiper belt (which has a very low mass and is indeed dynamically cold) but not the asteroid belt (which is dynamically hot). Interestingly, it is the asteroid belt, not the planetary system, that provides the most stringent constraint against this pile-up model.

Given our result, the Grand Tack model stands as the only current good explanation for both Mars and the asteroid belt. Of course, this does not mean that the Grand Tack is correct. In fact, we acknowledge that there is a difficult synchronism in the growth/migration histories of Jupiter and Saturn that is needed for the Grand Tack to work (see Raymond & Morbidelli 2014). Nevertheless there are as yet no satisfactory alternative models.

To summarize, we have shown that the simplest solution to the Mars problem – a mass deficit beyond 1–1.5 au – creates a new problem by underexciting the asteroid belt. The formation of Mars and the current architecture of the asteroid belt are deeply connected constraints for models of Solar system formation (Izidoro et al. 2014).

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